

## REGULATORY FACTORS OF SOIL CH<sub>4</sub> FLUXES IN DIFFERENT AGES OF OIL PALM ON TROPICAL PEATLAND, SARAWAK, MALAYSIA

Lulie Melling<sup>1</sup>, Kah Joo Goh<sup>2</sup>, Auldry Chaddy<sup>1</sup> and Ryusuke Hatano<sup>3</sup>

<sup>1</sup>*Tropical Peat Research Laboratory Unit (Chief Minister's Department), Jalan Badruddin, 93400 Kuching, Sarawak, Malaysia, +6082-241190, lulie\_melling@yahoo.com*

<sup>2</sup>*Advanced Agriecological Research Sdn. Bhd., Selangor, Malaysia*

<sup>3</sup>*Graduate School of Agriculture, Hokkaido University, Japan*

### SUMMARY

The objectives of this study were to quantify soil methane (CH<sub>4</sub>) fluxes and determine the regulatory factors affecting them in different ages of oil palm on tropical peatland in Sarawak, Malaysia. Monthly soil CH<sub>4</sub> fluxes were measured using a closed chamber method. Soil CH<sub>4</sub> fluxes decreased with palm age with values of 53.8, 23.1 and 17.5 µg C m<sup>-2</sup> h<sup>-1</sup> for 1, 5 and 7 year old palms, respectively. This may due to the release of trapped CH<sub>4</sub> in the peat vacuoles upon drainage and the increasing degree of CH<sub>4</sub> oxidation in the drained anoxic peat layers of older plantings. The significant correlations between rainfall, water table and water-filled pore space (WFPS) with soil CH<sub>4</sub> fluxes emphasized the importance of moisture in soil CH<sub>4</sub> fluxes in tropical peatland.

KEY WORDS: Oil palm, tropical peatland, soil CH<sub>4</sub> flux

### INTRODUCTION

Reclamation of tropical peatland for any agricultural activity, particularly oil palm cultivation, will require drainage and compaction to reduce the excess water and create an anoxic zone for better root respiration and also to increase the soil bulk density. Drainage and compaction have been suggested to affect the soil methane (CH<sub>4</sub>) flux in tropical peatland (Couwenberg, 2009).

Soil CH<sub>4</sub> flux could also be controlled by soil temperature (Couwenberg, 2009), precipitation (Inubushi *et al.*, 2003) and water-filled pore space (%WFPS) (Melling *et al.*, 2005). While the quantification of soil CH<sub>4</sub> fluxes and the factors that control them are well understood in the case of both temperate and boreal peat, there is still a lack of such studies for tropical peatlands. Therefore, the aims of this study were to quantify soil CH<sub>4</sub> fluxes and determine the regulatory factors affecting them in different ages of oil palm on tropical peatland.

### MATERIAL AND METHODS

This study was conducted in an oil palm plantation at Mukah, Sarawak, Malaysia (2° 51'N, 112° 13'E). The location has an equatorial climate with temperatures between 30.6 to 33.7° C and annual rainfall of more than 3000 mm. Drainage, with controlled water management and compaction, had been carried out during the establishment of the oil palm plantation. Three

sites, with oil palms of 1, 5 and 7 years old respectively, were selected. The peat soils at the site with 1 year old palms have been classified as *Typic Haplofibrists* whereas the sites with 5 and 7 year old palm had *Typic Haplohemists* peat soils. Detailed descriptions of each site are shown in Table 1.

Measurements of monthly soil CH<sub>4</sub> fluxes were conducted using a closed-chamber method from July 2006 until June 2008. Three random plots were set-up in each site for sampling of soil CH<sub>4</sub> gas. Environmental variables such as air temperature, relative humidity (RH) and soil temperature at 5 and 10 cm deep, depth of water table and rainfall were measured simultaneously with soil CH<sub>4</sub> flux measurement. The gas concentrations were determined in the laboratory using a gas chromatograph (Agilent 7890A).

Core soil samples were taken after gas sampling at each chamber point to determine soil bulk density and water-filled pore space (%WFPS). Core sample volumes were measured with a soil volume analyser (model Dik-1130, Daiki Rika Kogyo Co. Ltd) after which the samples were oven dried at 105 °C for 48 hours. The soil was also sampled using a peat auger in each chamber point at depths of 0-25 cm and 25-50 cm and analysed for selected chemical properties as shown in Table 1.

Cross correlation analysis was used to determine any correlation between soil CH<sub>4</sub> fluxes and rainfall, water table and WFPS. The significance of the cross correlation coefficient  $r$  at lag  $k$  can be assessed by the approximate test

$$t = r_k \sqrt{\frac{n-2}{1-r_k^2}},$$

with  $(n-2)$  degrees of freedom. A significance level of 0.05 was used for statistical significance, and a significance level of 0.10 was used for weak statistical significance in this study.

Table 1. Site description of oil palm plantation.

Age of palm (year)	1	5	7
	2° 50'N, 112° 13' E	2° 51'N, 112° 13'E	2° 53'N, 112° 12'E
Peat thickness (cm)	520	475	375
Bulk density (g cm <sup>-3</sup> )	0.23	0.21	0.22
WFPS* (%)	74.2	81.2	78.5
Water table (cm)	-56.4	-66.6	-55.6
Soil pH	3.6	3.8	3.9
Loss on ignition (%)	94.4	93.9	94.2
Total C (%)	55.6	55	56.1
Total N (%)	2.1	1.9	1.9
CEC* (cmol kg <sup>-1</sup> )	45.4	42.5	41.8
Base saturation (%)	40.7	60.4	61.8

## RESULTS

Mean monthly soil CH<sub>4</sub> fluxes decreased with palm age, with values of 53.8, 23.1 and 17.5  $\mu\text{g C m}^{-2} \text{ h}^{-1}$  for sites with 1, 5 and 7 year old palms, respectively. The corresponding cumulative soil CH<sub>4</sub> fluxes for the three sites were 465.5, 203.3 and 168.0  $\text{mg C m}^{-2} \text{ yr}^{-1}$  (Table 2).

Table 2. Mean monthly soil CH<sub>4</sub> flux and cumulative flux for oil palm plantation.

Age of palm (year)	Mean soil CH <sub>4</sub> flux ( $\mu\text{g C m}^{-2} \text{ h}^{-1}$ )	Cumulative soil CH <sub>4</sub> flux ( $\text{mg C m}^{-2} \text{ yr}^{-1}$ )
1	53.8	465.5
5	23.1	203.3
7	17.5	168.0

Soil CH<sub>4</sub> fluxes and rainfall showed positive correlations for all lag months in all sites. The highest correlation was recorded in 1 lagged month at the site with 1 year old palms but 2 lagged months in the sites with the older plantings. For the site with 1 year old palms, the correlation between soil CH<sub>4</sub> fluxes and rainfall at 1 lagged month was significant ( $p = 0.02$ ) whereas weakly significant correlations were observed in the other sites ( $p < 0.10$ ) (Fig. 1a). Depth of water table and soil CH<sub>4</sub> fluxes were found to be negatively correlated for all lags in all sites. Significant ( $p < 0.05$ ) peaks were observed in each site but in different lagged month (Fig. 1b). The correlation between soil CH<sub>4</sub> fluxes and WFPS were generally positive except in the case of the 1 year old palms, where they were inconsistent. Significant correlations ( $p < 0.05$ ) occurred in each site whereas a weakly significant correlation ( $p = 0.07$ ) was found at 0 lagged month at the site with 5 year old palms (Fig. 1c).

## DISCUSSION

The conversion of tropical peatland to oil palm initially increases soil CH<sub>4</sub> emissions probably due to the release of CH<sub>4</sub> trapped in the peat vacuoles prior to drainage. Furthermore, the compaction of peat for the planting of oil palm would improve gaseous diffusion. The decline in soil CH<sub>4</sub> emission with palm age might be attributed to the oxidation of soil CH<sub>4</sub> in the upper drained layers of peat and the decreasing production of CH<sub>4</sub> upon drainage. Higher rainfall enhances soil CH<sub>4</sub> flux (Inubushi *et al.*, 2003) by increasing the water table and WFPS, which in turn improves methanogenic activity. The positive correlations found between soil CH<sub>4</sub> fluxes and rainfall, water table and WFPS bear this out particularly in older palms (Fig. 1). Relatively constant monthly air and soil temperatures throughout the year may probably cause the lack of effect between soil CH<sub>4</sub> fluxes and soil temperatures for all palm ages (Davidson *et al.*, 2000). However, the indirect effect of temperature on soil CH<sub>4</sub> fluxes was a plausible reason for the inconsistent correlation between soil CH<sub>4</sub> fluxes and WFPS in 1 year old palms due to the disruption of methanogens activities in unshaded sunlight on soil

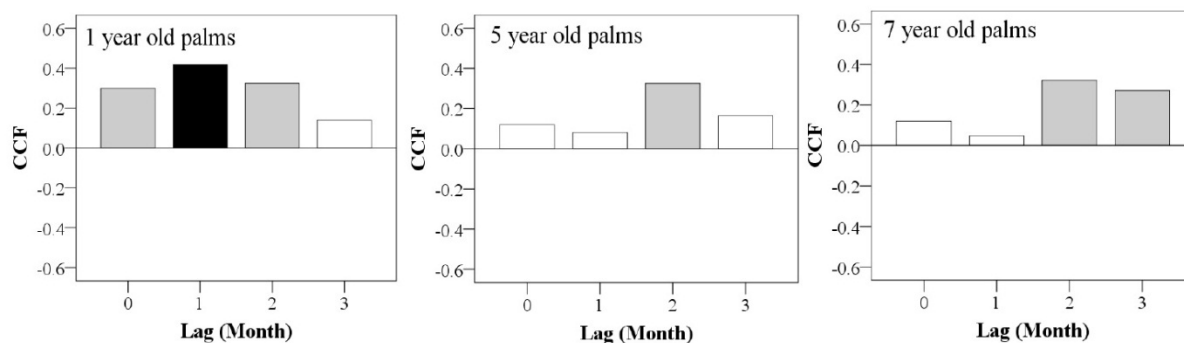
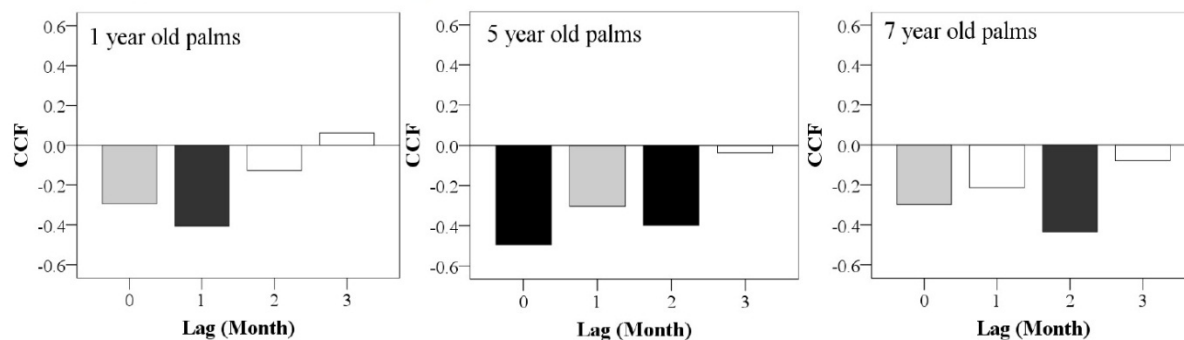
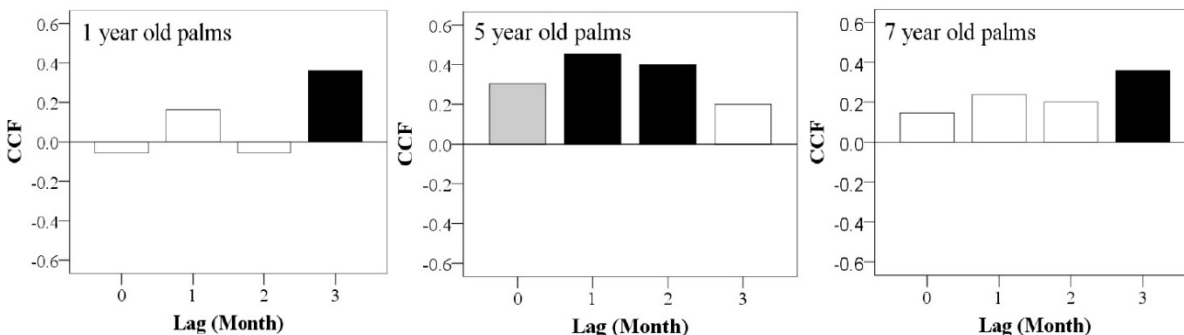
a) Soil CH<sub>4</sub> fluxes with rainfallb) Soil CH<sub>4</sub> fluxes with water table depthc) Soil CH<sub>4</sub> fluxes with WFPS

Fig. 1. Cross-correlation functions (CCF) for (a) soil CH<sub>4</sub> fluxes and rainfall, (b) soil CH<sub>4</sub> fluxes and water table depth and (c) soil CH<sub>4</sub> fluxes and water-filled pore space (WFPS) in oil palm plantation. Statistically significant cross-correlation is shown by black bars, weakly significant cross-correlation is shown by grey bars and non-significant cross-correlation is shown by white bars.

surface (Fig. 1c). Hence, higher temperature at the site with the youngest palms may produce high soil CH<sub>4</sub> fluxes through rapid diffusion rate but the effect is hampered by the loss of soil moisture. Soil CH<sub>4</sub> fluxes and cumulative soil CH<sub>4</sub> fluxes in this study were found to be relatively low compared with those reported by Inubushi *et al.* (2003) and Jauhainen *et al.* (2005) probably due to the higher water tables in their studies.

## CONCLUSION

This study clearly showed the importance of soil moisture in soil CH<sub>4</sub> fluxes of tropical peatland under oil palm. Further investigation and attention are imperative for better understanding and accurate interpretation of soil CH<sub>4</sub> flux variability in cultivated tropical peatland.

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## REFERENCES

- Couwenberg, J. (2009). *Emission factors for managed peat soils (organic soils, histosols) an analysis of IPCC default values*. Report produced for the UN-FCCC meetings in Bonn, June 2009.
- Davidson, E. A., Verchot, L. V., Cattanio, H., Ackerman, I. L. and Carvalho, J. E. M. (2000). Effects of soil water content on soil respiration in forests and cattle pastures of eastern Amazonia. *Biogeochemistry* **48**: 53–69.
- Inubushi, K., Furukawa, Y., Hadi, A., Purnomo, E. and Tsuruta, H. (2003). Seasonal changes of CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O fluxes in relation to land-use change in tropical peatlands located in coastal area of South Kalimantan. *Chemosphere* **52**: 603–608.
- Jauhiainen, J., Takahashi, H., Heikkinen, J.E.P., Martikainen, P.J. and Vasander, H. (2005). Carbon fluxes from a tropical peat swamp forest floor. *Global Change Biology* **11**: 1788–1797.
- Melling, L., Hatano, R. and Goh, K.J. (2005). Methane fluxes from three ecosystems in tropical peatland of Sarawak, Malaysia. *Soil Biology and Biochemistry* **37**: 1445–1453.